

# Experiments on the MET OFFICE method of river runoffs implementation in NEMO

Clément Vic      Anne-Marie Tréguier      Henrick Berger

Laboratoire de Physique des Océans

August 17, 2012

## Contents

<b>1 Brief description of the method</b>	<b>2</b>
1.1 How it works ?	2
1.2 How to use it ?	3
1.2.1 Namelist parameters	3
1.2.2 Input files	4
<b>2 Comparison with the top grid cell implementation</b>	<b>4</b>
2.1 First experiment	6
2.2 Second experiment	6
2.3 Third experiment	7
2.3.1 Runoffs prescribed on a large area	8
2.3.2 Runoffs prescribed on a few grid cells	9
2.3.3 Comparison with the waterway prescription	10
2.4 Remark on the temperature	11
<b>3 What should we remember about this new method ?</b>	<b>13</b>
<b>References</b>	<b>14</b>
<b>A Namelist</b>	<b>14</b>

This report is about the methods for prescribing river runoff in the Nucleus for European Modelling of the Ocean<sup>1</sup> (NEMO, Madec and the NEMO team, 2012) model.

In the NEMO ocean model as in most global circulation models, river runoffs used to be inserted on the top grid cells for precipitation. To better represent the reality of rivers where the fresh water flux has a nonzero depth, the vertical mixing is usually increased at the river mouth vicinity. From version 3.3 onwards it is possible to add river runoffs through many ocean levels. This new method has been coded and tested by R. Furner (U.K. MET OFFICE). In this report, we will present how this new method is implemented and how to use it, then we will compare it with the classical top grid cell implementation method. The configuration on which is based the comparison has been developed by Vic et al. (2012) in the framework for studying the dynamics of the Congo river plume in the Gulf of Guinea. We do not deliberately detail the parameters of this configuration more the dynamics of the plume to concentrate on the numerical aspects. Please refer to Vic et al. (2012) and to the namelist of the configuration in appendix A to get more details on the configuration and the dynamics.

## 1 Brief description of the method

### 1.1 How it works ?

The MET OFFICE method of runoff implementation differs from the classical surface-prescribed runoff method by the fact that freshwater is prescribed on several layers on the vertical. This results in adding a mass of freshwater (zero-salinity by default but it can be parameterized) in each grid cell concerned by the runoff prescription. According to routine `sbc_rnf_div`, it enters the mass conservation equation this way :

$$\frac{\partial}{\partial t} \vec{\nabla} \cdot \vec{V} = - \frac{1}{\rho_{water} h_{RNF}} \frac{RNF + RNF_b}{2}$$

where  $\vec{V}$  is the velocity,  $\rho_{water}$  the reference volumic mass of freshwater,  $h_{RNF}$  the depth on which runoffs are prescribed and  $RNF$  (and  $RNF\_b$ ,  $b$  for *before*) is the mass of freshwater entering the grid cell section every second :

$$RNF = \frac{\rho_{water} Q}{\sum_{i=1}^n S_{grid\ cell\ i}} \text{ (kg/m}^2/\text{s)}$$

where  $n$  is the number of grid cells where the runoffs are prescribed and  $S_{grid\ cell\ i}$  is the surface of the  $i^{th}$  grid cell.

According to the first equation, the same quantity of freshwater enters each grid cell along the vertical, generating a pressure gradient and horizontal velocities at every level as illustrated in figure 1 (this certainly consists in the major difference with the surface prescription where

---

<sup>1</sup><http://www.nemo-ocean.eu>

pressure gradient is only induced in the surface layer... We will confirm this in carrying out experiments later). Because of the varying size of the grid cells along the vertical, the pressure gradient is most important in the surface layers and velocities generated too.

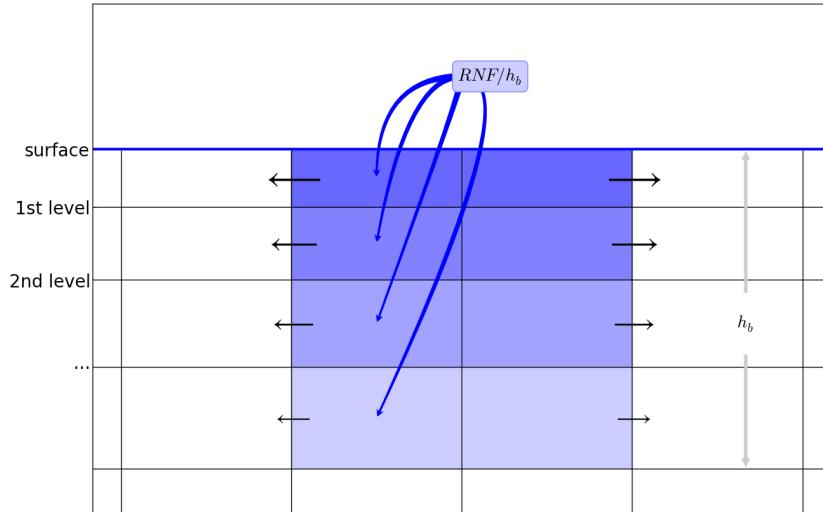


Figure 1: Schematic of the MET OFFICE method of multi-layer runoffs prescription. As suggested in the varying length of the arrows and in the gradient blue color of the cells, pressure induced by the freshwater flux and resulting velocities are more important at the surface.

## 1.2 How to use it ?

### 1.2.1 Namelist parameters

To prescribe the runoffs on multiple layers, one should set the namelist parameters this way (important parameters in red) :

```
!-----
&namsbc_rnf ! runoffs namelist surface boundary condition
!-----
!           ! file name ! frequency (hours) ! variable ! time interp. ! clim ! 'yearly'/ ! weights ! rotation !
!           !           ! (if <0 months) !   name   ! (logical) ! (T/F) ! 'monthly' ! filename ! pairing !
sn_rnf     = 'runoff', -1, 'sorunoff', .true., .true., 'yearly', '', '', ''
sn_cnf     = 'runoff', 0, 'socoeffr', .false., .true., 'yearly', '', '', ''
sn_s_rnf   = 'runoffs', -1, 'rosaline', .true., .true., 'yearly', '', '', ''
sn_t_rnf   = 'runoffs', -1, 'rottemper', .true., .true., 'yearly', '', '', ''
sn_dep_rnf = 'runoffs', 0, 'rodepth', .false., .true., 'yearly', '', '', ''

cn_dir      = './'      ! root directory for the location of the runoff files
ln_rnf_emp  = .false.  ! runoffs included into precipitation field (T) or into a file (F)
ln_rnf_mouth = .false. ! specific treatment at rivers mouths
rn_hrnf    = 10.e0    ! depth over which enhanced vertical mixing is used
rn_avt_rnf = 1.e-4    ! value of the additional vertical mixing coef. [m2/s]
rn_rfact   = 1.e0    ! multiplicative factor for runoff
ln_rnf_depth = .true. ! read in depth information for runoff
ln_rnf_tem  = .true.  ! read in temperature information for runoff
```

```
ln_rnf_sal = .true.      ! read in salinity information for runoff
```

The specific treatment at rivers mouths must not be activated, because it could increase the vertical mixing (vertical mixing coefficient `rn_avt_rnf` to a depth of `rn_hrnf`) on layers where the fresh water flux is already prescribed. When using simultaneously `ln_rnf_mouth = .true.` and one of `ln_rnf_depth/ln_rnf_tem/ln_rnf_sal = .true.`, one will get a warning in the `ocean.output`.

### 1.2.2 Input files

The method requires 4 input variables basically divided into 2 files :

- `runoff.nc` contains the variable `sorunoff` which contains the mass of fresh water per unity of surface per second at each point of the domain grid (`sorunoff` in kg/m<sup>2</sup>/s). It can have a time-dependency in hours (or it can be specified monthly if the parameter is negative), specified in the namelist.
- `runoffs.nc` contains the variables :
  1. `rosaline` is the salinity of the river runoff at each grid point. It is constant along the depth but can vary horizontally. If not specified, salinity is assumed to be fresh water (0 PSU).
  2. `rottemper` is the temperature of the river runoff at each grid point. It is constant along the depth but can vary horizontally. If not specified, temperature takes the SST value at each grid point along the depth.
  3. `rodepth` is the depth on which river runoff is prescribed. For the computation, it is adjusted to the bottom of the relevant grid box.

## 2 Comparison with the top grid cell implementation

To investigate the differences between the two methods, we use a 1/48° resolution idealized configuration based on the Congo river parameters (Vic et al., 2012); the namelist is provided in appendix A. The domain and bathymetry are shown on figure 2. The ocean is at rest and presents a constant salinity at each point of 35.5 PSU, and a homogenous profile of temperature representative of the Congo estuary vicinity (based on Levitus data). There is no forcing except the Congo runoff (no wind stress, no boundary flux, ...). The rate of flow  $Q$  is taken constant of 40000 m<sup>3</sup>/s (mean climatological value). Runoffs RNF are prescribed on the half upstream part of the waterway (cf figure 3) and its value is constant on the domain (see section 1 to calculate RNF).

We have carried out three basic experiments to compare the two methods :

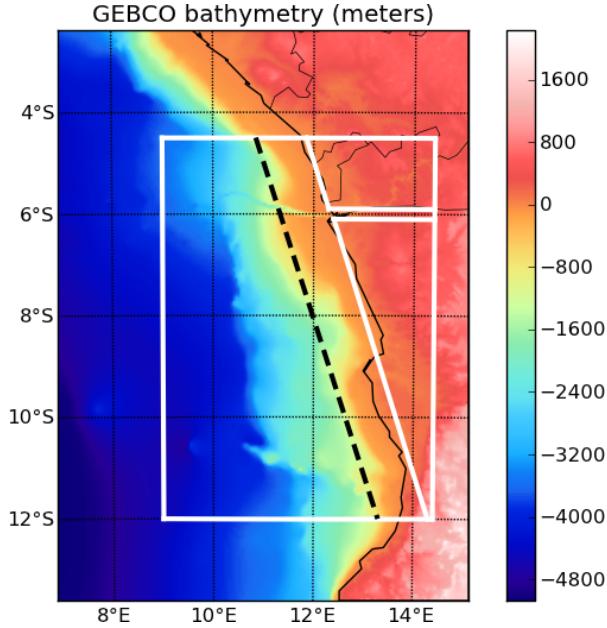


Figure 2: Domain and bathymetry of the idealized configuration. The Congo river estuary is represented as a zonal waterway at 6°S. The bathymetry is extremely simplified and has two levels : the waterway is at 30m depth and the ocean domain is at 2000m.

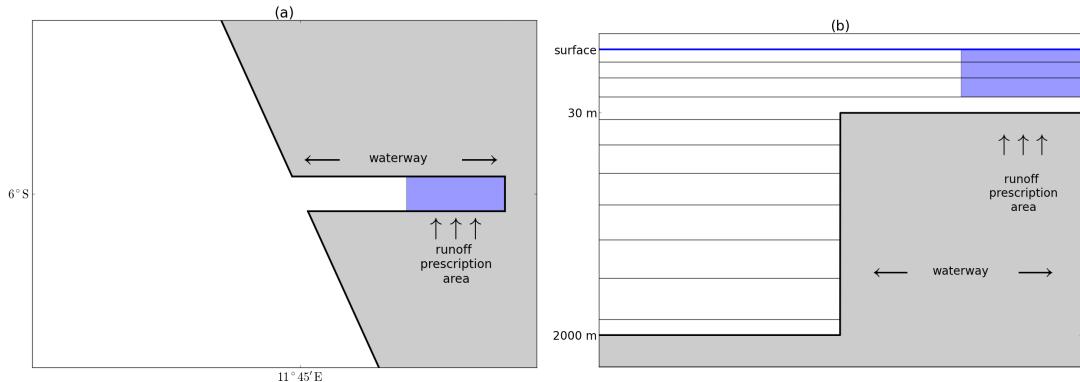


Figure 3: Schematic of the waterway and the runoff prescription area for (a) a plan view and (b) a section view. The MET OFFICE method allows to select several vertical levels. The waterway is 3 grid points wide in the prescription area and widens to 9 grid points at the river mouth to better represent the Congo mouth, see [Vic et al. \(2012\)](#) for further information.

1. First experiment : runoffs prescribed on the first layer in the MET OFFICE method must be compared with runoffs prescribed at the surface with the classical method. Both should give a thin layer of freshwater lying on the lower part of the ocean at rest.
2. Second experiment : runoffs prescribed on the first 10m depth in the MET OFFICE method must be compared with runoffs prescribed at the surface with increased vertical mixing  $K_z = 1\text{m}^2/\text{s}$  (value used to simulate convection) on 10m. We expect the two methods to

give a uniform stratification on the prescribed area.

3. Third experiment : to be closest to realistic configurations where there is usually no waterway, runoffs are prescribed on a configuration with a straight coastline and no waterway, on a more or less symmetrical area around the coordinates of the river mouth.

## 2.1 First experiment

The first experiment confirms that the two methods are exactly the same in the case of a surface prescription for the multi-layer method. We do not show the figures because the salinity and velocity fields are namely the same.

We now expect the two methods to give different results for a deeper prescription.

## 2.2 Second experiment

The multi-layer prescription brings freshwater directly in the subsurface layers whereas in the surface layer prescription, increased vertical mixing coefficient tends to freshen these layers.

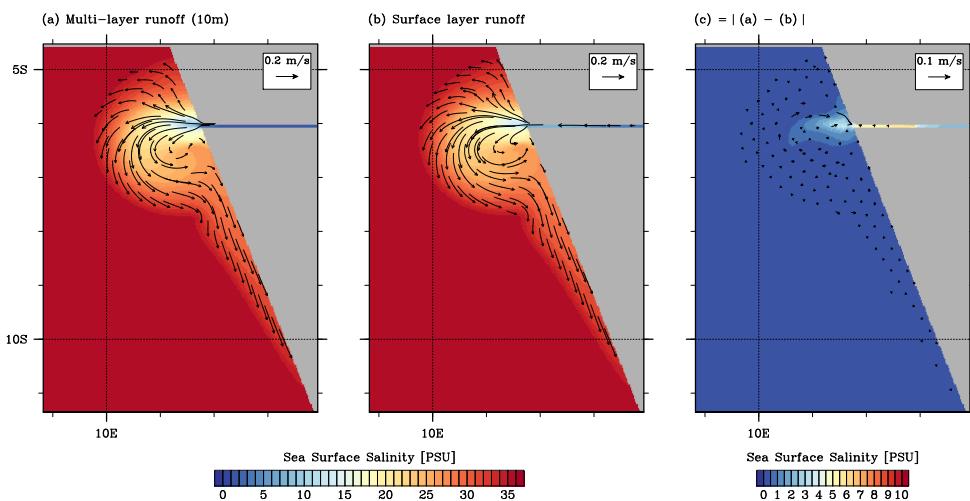


Figure 4: Sea surface salinity and surface velocity at 30 days for (a) runoffs prescribed on the first 10m (multiple ocean layers) and (b) runoffs prescribed at the surface and vertical mixing  $K_z$  is increased dramatically to  $1\text{m}^2/\text{s}$  in the first 10m. (c) is the difference between frame (a) and (b).

Figure 4 shows the sea surface salinity (SSS) and surface velocity for the two simulations at 30 days. Even if the SSS in the estuary reaches smaller values when runoff are prescribed on multiple layers (minimum of 1.0 PSU) than in the other case (minimum of 3.2 PSU), it does not seem to affect the velocity field in the ocean. Both methods show a classical anticyclonic circulation and a steady coastal current in the way of propagation of Kelvin waves (leaving the coast to the left because the experiment takes place in the southern hemisphere).

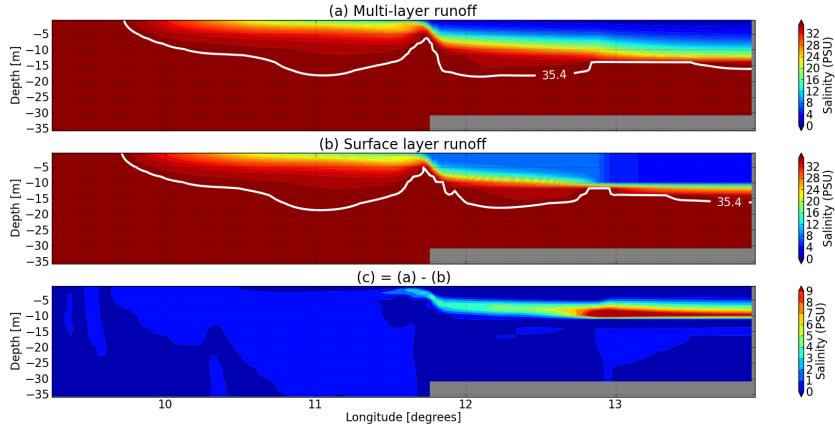


Figure 5: Salinity sections along the waterway at 30 days for (a) runoffs prescribed on the first 10m (multiple ocean layers) and (b) runoffs prescribed at the surface and vertical mixing  $K_z$  is increased dramatically to  $1\text{m}^2/\text{s}$  in the first 10m. The white line represents isohaline 35.4 PSU which is the limit of influence of fresh water. (c) is the difference between frame (a) and (b).

Figure 5 shows the salinity sections through the waterway at 30 days of simulation. The major difference concerns the area where the runoffs are prescribed : the increased vertical mixing (panel (b)) seems to be more «powerful» in homogenizing the water column on the first 10m than the runoffs prescription on multiple layers. The salinity increases smoothly in the multi-layer prescription, which is maybe more representative of the real stratification of estuaries. The salinity profiles in the basin do not present any striking difference with each other. The influence of freshwater is represented by isohaline 35.4 PSU, and has the same shape in both cases. We could explain this by the fact that in the downstream part of the waterway the buoyancy adjustment is made rapidly : fresher waters reach the surface and the more common vertical diffusion ( $\propto 10^{-5}\text{ m}^2/\text{s}$ ) tends to stabilize the flux. In the end, the difference between the two salinity fields vanishes along the waterway (panel (c)) and the salinity profiles are equivalent at the river mouth.

Doing the same experiment with a more classical vertical diffusion coefficient ( $10^{-4}\text{ m}^2/\text{s}$ ) at the river mouth (specific treatment) for the surface prescription gives similar results. Furthermore, it seems to be possible to get the same stratification and velocity fields in adjusting this coefficient.

### 2.3 Third experiment

The third experiment tests the methods in a more realistic case (without a waterway like in former academic shemes). The runoffs are prescribed on an area centered on the position of the river mouth ( $6^{\circ}00'\text{S}$  ;  $11^{\circ}45'\text{E}$ ), on 10m for the multi-layers method and at the surface method with increased  $K_z = 10^{-4}\text{ m}^2/\text{s}$  in the other method.

### 2.3.1 Runoffs prescribed on a large area

Like in former experiments, in order to keep the same order of magnitude of the RNF value and to generate comparable velocities, runoffs are prescribed on 101 grid cells.

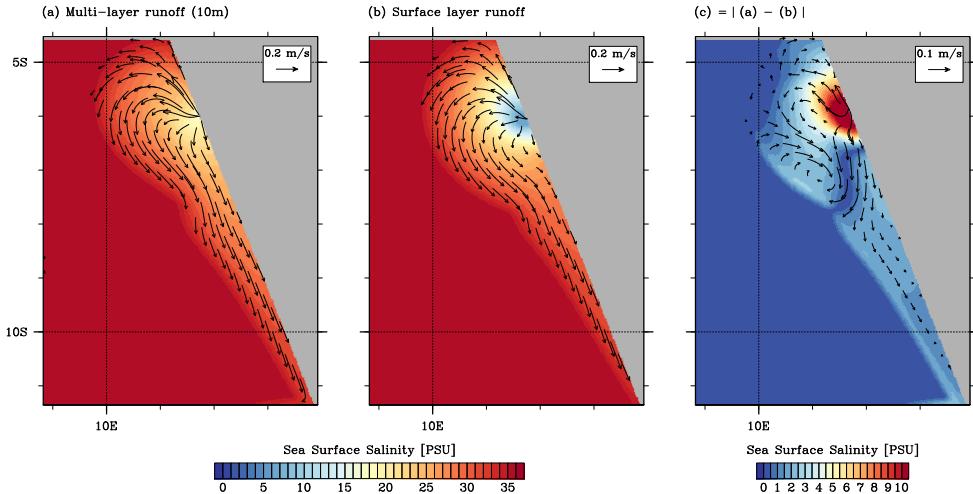


Figure 6: Sea surface salinity and surface velocity at 30 days for (a) runoffs prescribed on the first 10m (multiple ocean layers) and (b) runoffs prescribed at the surface and vertical mixing  $K_z$  is increased to  $10^{-4} \text{m}^2/\text{s}$  in the first 10m. (c) is the difference between frame (a) and (b).

The first comment that comes from the comparison of figures 4 and 6 is that the presence of the waterway does not affect dramatically the shape of the dynamics. We still have the anticyclonic circulation and the coastal current. Nonetheless, small differences concern the cross-shore extent compared with the width of the coastal current : on the one hand, the waterway controls the flux which suddenly reaches the basin at rest and preferentially goes offshore following its former direction and then is deflected to the left by the Coriolis force. On the other hand, with no waterway, the flux seems to spread more isotropically in all the directions (not preferentially to the offshore direction).

The most striking difference is about the SSS : the two methods present a difference which is maximum at the bulge core and reaches 10 PSU (panel (c) of figure 6). According to figure 7 differences are confined to the prescription area and in the first 3 m, then they vanish with the flow. The multi-layer prescription actually better distribute the amount of freshwater with depth and it results a less desalinated water at the surface.

The other major difference concerns the variation of velocity (and salinity) with the depth. It derives from the fact that the multi-layer prescription brings a mass of water in every layer of the column until 10m, so that a pressure gradient is created in each layer. As a consequence, velocities are properly induced in each layer and the vertical shear is less important than with the surface layer prescription : we can see it in figure 8 where the velocity field at 5m depth is less attenuated in the multi-layer prescription (panel (a)) than in the surface layer prescription

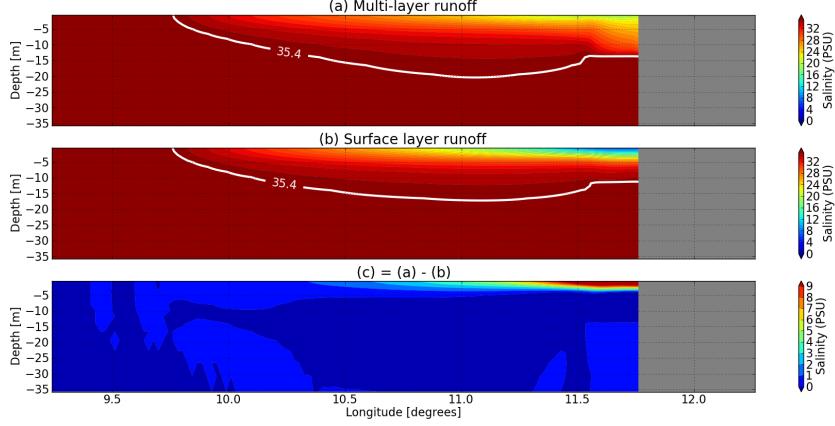


Figure 7: Salinity section at 6°S (runoffs are prescribed symmetrically along this axis) at 30 days for (a) runoffs prescribed on the first 10m (multiple ocean layers) and (b) runoffs prescribed at the surface and vertical mixing  $K_z$  is increased to  $10^{-4} \text{ m}^2/\text{s}$  in the first 10m. The white line represents isohaline 35.4 PSU which is the limit of influence of fresh water. (c) is the difference between frame (a) and (b).

(panel (b)), compared to the surface velocity field. The difference of velocity between the two methods even reaches 10 cm/s in the bulge circulation (panel (c)).

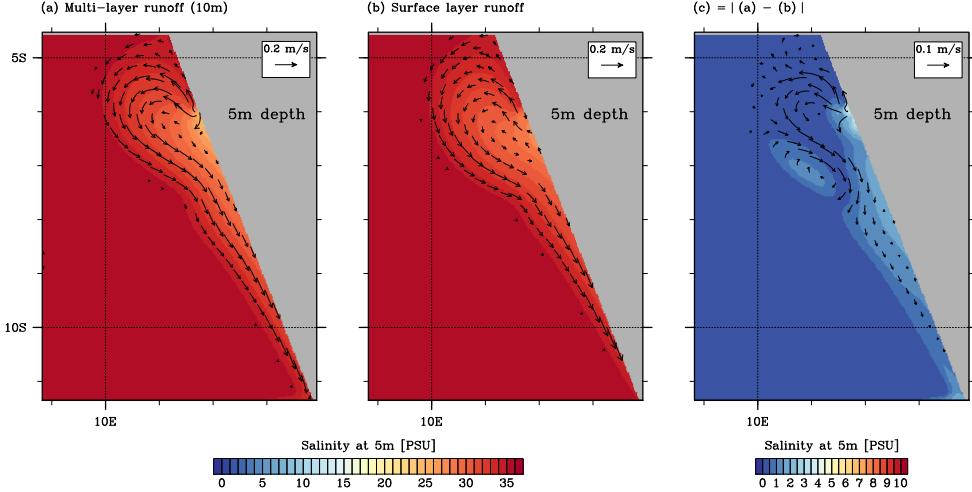


Figure 8: Salinity and velocity at 30 days at 5m depth for (a) runoffs prescribed on the first 10m (multiple ocean layers) and (b) runoffs prescribed at the surface and vertical mixing  $K_z$  is increased to  $10^{-4} \text{ m}^2/\text{s}$  in the first 10m. (c) is the difference between frame (a) and (b).

### 2.3.2 Runoffs prescribed on a few grid cells

In order to avoid prescription on ocean grid cells where there is no physical reason to bring freshwater (as we did in the former case with 101 grid cells affected by the prescription<sup>2</sup>),

<sup>2</sup>which represent a surface of  $\approx 500 \text{ km}^2$

we confine the prescription on a small patch around the river mouth (here we take 6 points). Therefore, the value of RNF is increasing by two orders of magnitude ( $\approx 1 \text{ kg/m}^2/\text{s}$ ). Simulation with surface prescription aborted quickly as salinity could not be evacuated in the deeper layers and reached negative values.

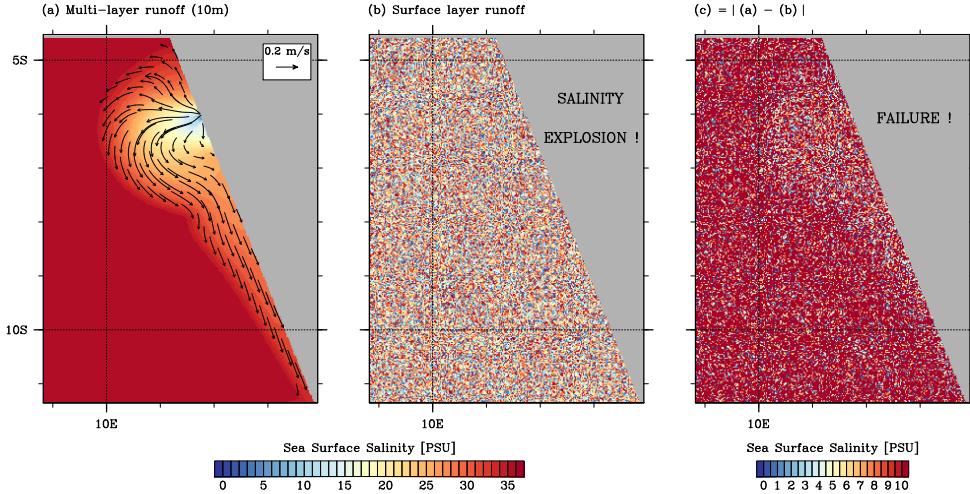


Figure 9: Sea surface salinity and surface velocity at 30 days for (a) runoffs prescribed on the first 10m (multiple ocean layers) and (b) runoffs prescribed at the surface and vertical mixing  $K_z$  is increased to  $10^{-4} \text{ m}^2/\text{s}$  in the first 10m. (c) is the difference between frame (a) and (b). In case (b), the simulation aborts quickly as salinity becomes negative.

Figure 9 must be compared with figures 6 and 4. The dynamics of the surface layer seems to be closer to the one generated with the waterway than to the one without waterway with prescription on a large area. In fact, the anticyclonic bulge is less distorted to the coast and velocities are more comparable with a maximum value of 1.4 m/s at the river mouth and velocities in the bulge of 25 cm/s and of 20 cm/s in the coastal current. Moreover, the minimum of salinity reached at the river mouth is closer ( $\approx 9 \text{ PSU}$ ).

Salinity section in this case (Figure 10) differs from the usual ones (*e.g.* figure 7) by the depth reached by fresh water. The depth of isohaline 35.4 PSU only reaches 15 m whereas it goes deeper than 20 m in the waterwayless reference case.

### 2.3.3 Comparison with the waterway prescription

The two experiments carried out in the waterwayless cases must be compared with the waterway case in order to observe which type of patch ( $\approx 100$  points-prescription one or  $\approx 1$  point-prescription one) best fits with the reality. As we described sooner, the shape of the plume is more realistic in the 6-point-prescription experiment than in the large patch prescription one as illustrated in figure 11.

In terms of dynamics and salinity repartition, differences are minimum with the small patch prescription (panels (c) and (f)). The major differences concern the recirculation formed just

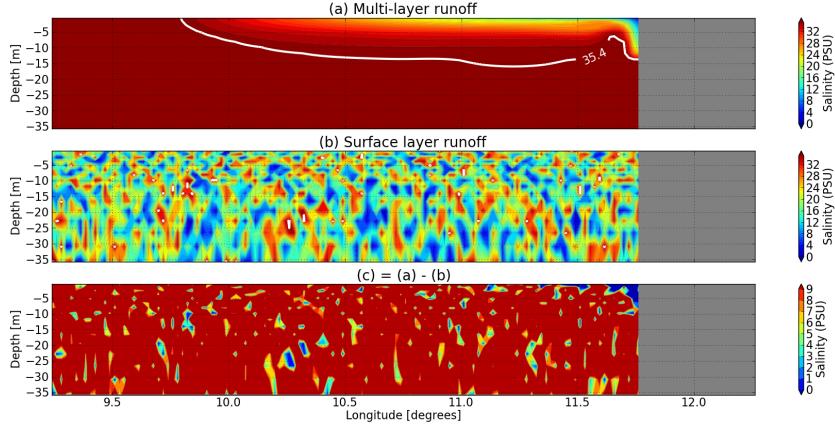


Figure 10: Salinity section at 6°S at 30 days for (a) runoffs prescribed on the first 10m (multiple ocean layers) and (b) runoffs prescribed at the surface and vertical mixing  $K_z$  is increased to  $10^{-4} \text{ m}^2/\text{s}$  in the first 10m. The white line represents isohaline 35.4 PSU which is the limit of influence of fresh water. (c) is the difference between frame (a) and (b). In case (b), the simulation aborts quickly as salinity becomes negative.

below the river mouth : it is more pronounced in the waterway prescription than in the patches ones. Nevertheless, it affects only the vicinity of the river mouth and differences of velocity vanish as we reach the frontier of the bulge.

Considering the stratification, it seems to be qualitatively the same, and it is difficult to say which one is better because it certainly depends on the river represented, and a comparison should be done with the observations to determine which profile of salinity best fits.

## 2.4 Remark on the temperature

We also did experiments on temperature but without precise sensibility tests. The effect observed is that decreasing the freshwater temperature (as it happens in reality where freshwater is often colder than the ocean) increases its density and tends to reduce the density gradient and the velocities induced.

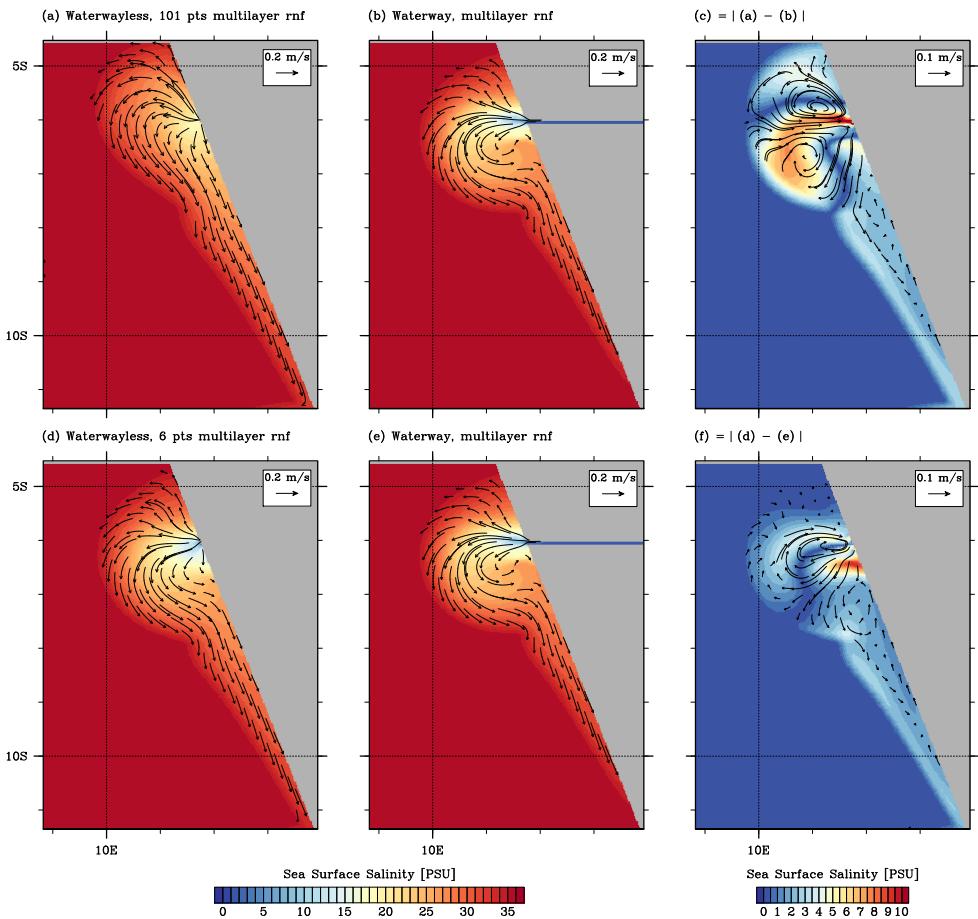


Figure 11: Salinity and velocity at 30 days at 5m depth for runoffs prescribed on multilayer (10 m) : without waterway on 101 grid cells (a) and 6 grid cells (d), with waterway (b) $\equiv$ (e). (c) is the difference between frames (a) and (b). (f) is the difference between frames (d) and (e).

### 3 What should we remember about this new method ?

The MET OFFICE method aims to prescribe runoffs on several vertical levels to be more realistic concerning rivers outflows as they usually enter ocean with a non-zero depth.

- The main difference with the surface prescription is that a pressure gradient is involved in each layer so that velocities are generated at each level.
- Concerning the stratification, one can obtain the same values of salinity along with the depth in the surface prescription method by tuning vertical mixing coefficient (from  $10^{-4}$  m<sup>2</sup>/s as with the original specific treatment at river mouth, to 1 m<sup>2</sup>/s to get a constant salinity along with the depth). Nonetheless, even if the result is the same for the salinity stratification, the fact that we do not need to increase dramatically Kz is a gain in terms of representation of a more realistic physics.
- The most important breakthrough is that runoffs can be prescribed on a minimum number of grid cells, as freshwater is evacuated easier. As a result, one can divide by two orders of magnitude the number of grid cells in comparison with the surface prescription. It enables not to enhance vertical mixing as it was the case but for no other physical reason than homogenizing freshwater flux along with the depth, so that we can say that the new method is more physical. Furthermore, we obtain better results in terms of dynamics with smaller patches.

## References

Madec, G. and the NEMO team (2012). *NEMO ocean engine*, volume 3.4. Pôle de modélisation de l’Institut Pierre-Simon Laplace.

Vic, C., Treguier, A.-M., and Berger, H. (2012). Dynamique d’un panache fluvial en zone équatoriale, modélisation numérique et application au cas du Congo. Master’s thesis, ENSTA ParisTech and Paris VI UPMC.

## A Namelist



```

ln_tsd_init = .false. ! Initialisation of ocean T & S with T & S input data (T) or not (F)
ln_tsd_trdmp = .false. ! damping of ocean T & S toward T & S input data (T) or not (F)
/
!=====
!*** Surface Boundary Condition namelists ****
!=====
!&namsbc ! surface boundary condition
!&namsbc_ana ! analytical formulation
!&namsbc_rnf ! flux formulation
!&namsbc_fix ! Clio bulk formulae formulation
!&namsbc_core ! CORE bulk formulae formulation
!&namsbc_mfs ! MFS bulk formulae formulation
!&namsbc_cpl ! CouPLed formulation
!&namsbc_qsr ! penetrative solar radiation
!&namsbc_lrf ! river runoffs
!&namsbc_apr ! Atmospheric Pressure
!&namsbc_ssr ! sea surface restoring term (for T and/or S)
!&namsbc_alb ! albedo parameters
!=====

!&namsbc ! Surface Boundary Condition (surface module)
!=====
mn_fsbcc = 1 ! frequency of surface boundary condition computation
               ! (also = the frequency of sea ice model call)
ln_ana = .true. ! analytical formulation
ln_fix = .false. ! flux formulation
ln_bik_clio = .false. ! Clio bulk formulation
ln_bik_core = .false. ! CORE bulk formulation
ln_bik_mfs = .false. ! MFS bulk formulation
ln_cpl = .false. ! Coupled formulation
ln_apr_dyn = .false. ! Patm gradient added in ocean & ice Eqs.
mn_ice = 0 ! no ice boundary condition
           !=0 use observed ice-cover
           !=2 ice-model used
           != daily mean to diurnal cycle on short wave
           != runoff
           != Sea Surface Restoring on T and/or S
           != Freshwater Budget: =0 unchecked
               !=1 global mean of e-pr set to zero at each time step
               !=2 annual global mean of e-pr set to zero
               !=3 global emp set to zero and spread out over erp area
               != Neutral drag coefficient read from wave model (T => fill namsbc_wave)
ln_dmd2c = .false. ! analytical surface boundary condition
ln_rnf = .true. ! gently increase the stress over the first ntau_rst time-steps
ln_ssr = .false. ! uniform value for the i-stress
mn_fvb = 0 ! uniform value for the j-stress
ln_cdgw = .false. ! uniform value for the solar radiation
/
!&namsbc_ana ! analytical surface boundary condition
!=====
mn_tau000 = 0 ! uniform value for the total heat flux
rn_utau0 = 0. ! uniform value for the solar radiation
rn_qns0 = 0. ! uniform value for the total heat flux
rn_asf0 = 0. ! uniform value for the solar radiation
/
!&namsbc_flx ! surface boundary condition : flux formulation
!=====
!&namsbc ! file name : frequency (hours) : variable : time interp.
!&namsbc ! (if <0 months) : name : (Logical)
sn_utau = 'utau' , 24 , 'utau' , .false.
sn_vtau = 'vtau' , 24 , 'vtau' , .false.
sn_qtau = 'qtau' , 24 , 'qtau' , .false.
sn_qsr = 'qsr' , 24 , 'qsr' , .false.
sn_emp = 'emp' , 24 , 'emp' , .false.
cn_dir = './' ! root directory for the location of the flux f
/
!&namsbc_clio ! nambsc_clio Clio bulk formulæ
!=====
!&namsbc ! file name : frequency (hours) : variable : time interp.
!&namsbc ! (if <0 months) : name : (Logical)
sn_utau = 'taux_1m' , -1 , 'soztaux' , .true.
sn_vtau = 'tavy_1m' , -1 , 'sometay' , .true.
sn_vndm = 'fix' , -1 , 'sociiov' , .true.
sn_tair = 'fix' , -1 , 'sociot2' , .true.
sn_humi = 'fix' , -1 , 'socioc' , .true.
sn_ccov = 'fix' , -1 , 'sociopl' , .true.
sn_prec = 'fix' , -1 , 'sociopl' , .true.
cn_dir = './' ! root directory for the location of the bulk f
/
!&namsbc_core ! nambsc_core CORE bulk formulæ
!=====
!&namsbc ! file name : frequency (hours) : variable : time interp.
!&namsbc ! (if <0 months) : name : (Logical)
sn_vndm = 'u_10-15JUNE2009_orca2' , 6 , 'U_10_M0D' , .false.
sn_vndj = 'v_10-15JUNE2009_orca2' , 6 , 'V_10_M0D' , .false.
sn_qsr = 'ncar_rad.15JUNE2009_orca2' , 24 , 'SDDN_M0D' , .false.
sn_qlw = 'ncar_rad.15JUNE2009_orca2' , 24 , 'LUDN_M0D' , .false.
sn_humi = 'q_10-15JUNE2009_orca2' , 6 , 'Q_10_M0D' , .false.
sn_prec = 'ncar_precip.15JUNE2009_orca2' , -1 , 'PRC_M0D1' , .false.
sn_snow = 'ncar_precip.15JUNE2009_orca2' , -1 , 'SNOW' , .false.
sn_tduf = 'taudif_core' , 24 , 'taudif' , .false.
cn_dir = './' ! root directory for the location of the bulk f
ln_2m = .false. ! air temperature and humidity referenced at 2m
ln_taudif = .false. ! HF tau contribution: use "mean of stress modis"
rn_pfac = 1. ! multiplicative factor for precipitation (total)
/
!&namsbc_mfs ! namsbc mfs MFS bulk formulæ
!=====

```



```

! namagrif : AGRIF zoom                                         ("key_agriff")
rn_cloud      = 0.06   ! cloud correction to snow and ice albedo
rn_albice     = 0.53   ! albedo of melting ice in the arctic and antarctic
rn_alphaPhi   = 0.80   ! coefficients for linear interpolation used to
rn_alphaPhi  = 0.65   ! compute albedo between two extremes values
rn_alphaPhi  = 0.72   ! (Pyne, 1972)
/
! namalbc      ! lateral boundary condition ***
! namalbc      ! lateral momentum boundary condition
! namcla       ! cross land advection
! namobc       ! open boundaries parameters
! namobc       ! ("key_ocbc")
! namagrif    ! agrif nested grid ( read by child model only )
! namagrif    ! ("key_agriff")
! namby        ! unstructured open boundaries
! namby        ! ("key_by_tides")
! namide       ! tidal forcing at open boundaries
! namide       ! ("key_by_tides")
/
! namalbc      ! lateral momentum boundary condition
! namalbc      ! lateral momentum boundary condition
!
! namalbc      ! lateral momentum boundary condition
!
rn_shlat     = 0.          ! shlat = 0 ! 0 < shlat < 2 ! 2 < shlat
                           ! free slip ! partial slip ! no slip ! strong slip
                           ! consistency of vorticity boundary condition with analytical eqs.
/
! namcla       ! cross land advection
! namcla       ! advection between 2 ocean pts separates by land
rn_cla       = 0.          ! advection between 2 ocean pts separates by land
/
! namobc       ! open boundaries parameters
! namobc       ! ("key_ocbc")
!
ln_ocbc_clim = .false.   ! climatological obc data files (T) or not (F)
ln_vol_cst   = .true.    ! impose the total volume conservation (T) or not (F)
ln_obe_fla   = .false.   ! flather open boundary condition
rn_obeData   = 1.         ! = 0 the obc data are equal to the initial state
                           ! = 1 the obc data are read in 'obc.dta' files
cn_obeData   = 'annual', ! set to annual if obc datafile hold 1 year of data
                           ! set to monthly if obc datafile hold 1 month of data
rn_dpsin     = 1.         ! damping time scale for inflow at east open boundary
rn_dpwin    = 1.         ! -           -           west   -
rn_dpinin   = 1.         ! -           -           north  -
rn_dpsin    = 1.         ! -           -           south  -
rn_dpeob    = 3000.      ! time relaxation (days) for the east open boundary
rn_dpob    = 15.          ! -           -           west   -
rn_dpob    = 3000.      ! -           -           north  -
rn_dpsob    = 15.          ! -           -           south  -
rn_volemp   = 1.          ! = 0 the total volume change with the surface flux (E-P-R)
                           ! = 1 the total volume remains constant
/
! namangrif   ! tide parameters (#ifdef key_tide)
!
! ln_tide_pot = use tidal potential forcing
! nb_harmo   = number of constituents used
! names(1)   = 'N2', 'K1', etc name of constituent
&nam_tide
ln_tide_pot = .true.
nb_harmo   = 11
cname(1)   = 'M2',
cname(2)   = 'S2',
cname(3)   = 'N2',
cname(4)   = 'K1',
cname(5)   = 'O1',
cname(6)   = 'Q1',
cname(7)   = 'M4',
cname(8)   = 'K2',
cname(9)   = 'P1',
cname(10)  = 'Mf',
cname(11)  = 'Mm',
/
! &namby      ! unstructured open boundaries
! &namby      ! ("key_bdy")
!
! nb_bdy = 1
! ln_coords_file = .true.
! ln_coords_file = 'coordinates.bdy.nc'
! ln_mask_file = .false.
! ln_mask_file = 'coordinates.bdy.nc'
! ln_mask_file = '',
! m_dyn2d    = 2
! m_dyn2d_dta = 3
!
! number of open boundary sets
! =T : read bdy coordinates from file
! =F : read mask from file
! name of mask file (if ln_mask_file=.TRUE.)
! boundary conditions for barotropic fields
! = 0, bdy data are equal to the initial state
! = 1, bdy data are read in 'bysize .nc' files
! = 2, use tidal harmonic forcing data from files
! = 3, use external data AND tidal harmonic forcing
! boundary conditions for T and S
! = 0, bdy data are equal to the initial state
! = 1, bdy data are read in 'bysize .nc' files
! width of the relaxation zone
! total volume correction (see mn_volct parameter)
! = 0, the total water flux across open boundaries is zero
/

```



```

rn_alpha      = 2.0e-4   ! = 2, linear; rho(T,S) = rho0 * ( rho * S - ralpha * T ) /
rn_beta       = 7.7e-4   ! thermal expansion coefficient (rn_eos=1 or 2)
rn_eos        = 2          ! saline expansion coefficient (rn_eos=2)

!-----!
!-----!> advection
!-----!
ln_tradv_cen2 = .false. ! 2nd order centered scheme
ln_tradv_tvrd = .true.  ! TVD scheme
ln_tradv_muscl = .false. ! MUSCL scheme
ln_tradv_muscl2 = .false. ! MUSCL2 scheme + ce2 at boundaries
ln_tradv_ubs  = .false. ! UBS scheme
ln_tradv_qck  = .false. ! QUICKST scheme
ln_tradv_ldf  ! lateral diffusion scheme for tracer
!-----!
!-----!> advection scheme for tracer
!-----!
ln_tradv_lap  = .false. ! 2nd order centered scheme
ln_tradv_tvrd = .true.  ! TVD scheme
ln_tradv_muscl = .false. ! MUSCL scheme
ln_tradv_ubs  = .false. ! UBS scheme
ln_tradv_qck  = .false. ! QUICKST scheme
ln_tradv_ldf  ! lateral diffusion scheme for tracers
!-----!
!-----!> Operator type:
!-----!
ln_traldf_level = .false. ! iso-level
ln_traldf_hor   = .true.  ! horizontal (geopotential) (needs "key_ldfslp" when ln_sco=T)
ln_traldf_iso   = .false. ! iso-neutral (needs "key_ldfslp")
!-----! Griffies parameters (all need "key_ldfslp")
ln_traldf_grif  = .false. ! use griffies triads
ln_traldf_gdia  = .false. ! output griffies eddy velocities
ln_traldf_iso   = .false. ! pure lateral mixing in ML
ln_bomix_grif   = .false. ! lateral mixing on bottom
!-----!
!-----! Coefficients
!-----!
! Eddy-induced (GM) advection always used with Griffies; otherwise needs "key_traldf_eiv"
! Value rn_aeiv_0 is ignored unless = 0 with Held-Larichek spatially varying aeiv
! (key_traldf_c2d & key_traldf_eiv & key_orca_r2, -r1 or -r05)
rn_aeiv_0       = 0.          ! eddy induced velocity coefficient [m2/s]
rn_drt_0        = -1.e9        ! horizontal eddy diffusivity for tracers [m2/s]
rn_abrb_0       = 0.          ! background eddy diffusivity for ldf_iso [m2/s]
!-----!
!-----!> tracer: T & S newtonian damping
!-----!
ln_tradmp      = .false. ! add a damping term (T) or not (F)
rn_hdfp        = -1          ! horizontal shape =-1, damping in Med and Red Seas only
                           ! =XX, damping poleward of XX degrees (XX>0)
                           ! + F (distance-to-coast) + Red and Ned Seas
mn_zdmp        = 1           ! vertical shape =0 damping throughout the water column
                           ! =1 no damping in the mixed layer (rho criteria)
                           ! =2 no damping in the mixed layer (rho criteria)
                           ! surface time scale of damping [days]
                           ! bottom time scale of damping [days]
rn_surf        = 50.         ! depth of transition between rn_surf and rn_bot [meters]
rn_dep         = 800.        ! create a damping_coeff NetCDF file (=1) or not (=0)
mn_file        =
!-----!
!-----!> Dynamics namelists ***
!-----!
!-----!> Dynamics namelists ***
!-----!
!-----!> namdyn_adv formulation of the momentum advection
!-----!
!-----!> namdyn_vor advection scheme
!-----!
!-----!> namdyn_hpg hydrostatic pressure gradient
!-----!
!-----!> namdyn_spf surface pressure gradient
!-----!
!-----!> namdyn_ldf lateral diffusion scheme
!-----!
!-----!> &namdyn_adv formulation of the momentum advection
!-----!
!-----!> In_dynadv_vec = .true. ! vector form (T) or flux form (F)
!-----!
!-----!> In_dynadv_cen2= .false. ! flux form - 2nd order centered scheme
!-----!
!-----!> In_dynadv_ubs = .false. ! flux form - 3rd order UBS scheme
!-----!
!-----!> &namdyn_vor option of physics/algorithm (not control by GPP keys)
!-----!
!-----!> In_dynvor_ene = .false. ! entrophy conserving scheme
!-----!
!-----!> In_dynvor_ens = .false. ! energy conserving scheme
!-----!
!-----!> In_dynvor_mix = .false. ! mixed scheme
!-----!
!-----!> In_dynvor_enen = .true. ! entropy & enstrophy scheme
!-----!
!-----!> &namdyn_hpg Hydrostatic pressure gradient option
!-----!
!-----!> In_hpg_zco = .false. ! z-coordinate - full steps
!-----!
!-----!> In_hpg_zps = .true. ! z-coordinate - partial steps (interpolation)
!-----!
!-----!> In_hpg_sco = .true. ! s-coordinate (standard Jacobian formulation)
!-----!
!-----!> In_hpg_djic = .false. ! s-coordinate (Hensley Jacobian with Cubic polynomial)
!-----!
!-----!> In_hpg_pj = .false. ! s-coordinate (Pressure Jacobian scheme)
!-----!
!-----!> In_dynam_hpg_implicit = .false. ! time stepping: semi-implicit time scheme (T)
!-----!> centered time scheme (F)
!-----!
!-----!> &namdyn_spf surface pressure gradient (GPP key only)
!-----!
!-----!> ! explicit free surface
!-----!
!-----!> ! filtered free surface
!-----!
!-----!> ! split-explicit free surface
!-----!
!-----!> &namdyn_ldf lateral diffusion on momentum
!-----!
!-----!> ! Type of the operator :
!-----!
!-----!> ln_dyndfd_lap = .false. ! laplacian operator
!-----!
!-----!> ln_dyndfd_bilap = .true. ! bilaplacian operator
!-----!
!-----!> mn_file : Direction of action :

```

```

ln_dyndf_level = .false. ! iso-level
ln_dyndf_hor = .true. ! horizontal (geopotential)
ln_dyndf_iso = .false. ! iso-neutral
! : Coefficient
rn_ahm_0_lap = 100000. ! horizontal laplacian eddy viscosity [m2/s]
rn_ahm_0_bdfp = 0. ! background eddy viscosity for ldf_1so [m2/s]
rn_ahm_0_bdfp = -1.e9 ! horizontal bilaplacian eddy viscosity [m4/s]
/
! ===== Tracers & Dynamics vertical physics name lists
! =====
namdf ! vertical physics
namdf_ric ! richardson number dependent vertical mixing ("key_zdfrc")
namdf_zke ! TKE dependent vertical mixing ("key_zdfzke")
namdf_zpp ! KPP dependent vertical mixing ("key_zdfzpp")
namdf_zdm ! double diffusive mixing parameterization ("key_zdfzdm")
namdf_tmx ! tidal mixing parameterization ("key_zdftmx")
!
! ===== vertical physics
namdf ! richardson number dependent vertical diffusion ("key_zdfrc")
rn_avm0 = 1.2e-4 ! vertical eddy viscosity [m2/s] (background Kz if not ("key_zdfcst"))
rn_av0 = 1.2e-5 ! vertical eddy diffusivity [m2/s] (background Kz if not ("key_zdfcst"))
rn_avb = 0 ! profile for background avt & avn (-1) or not (-0)
rn_havtb = 0 ! horizontal shape for avt (-1) or not (-0)
rn_evdm = 1 ! enhanced vertical diffusion (evd) (T) or not (F)
rn_avend = 100. ! evd apply on tracer (-0) or on tracer and momentum (=1)
rn_zdfnpc = .false. ! Non-Penetrative Convective algorithm (T) or not (F)
rn_npc = 1 ! frequency of application of npc
rn_nppc = 365 ! npc control print frequency
rn_zdfexp = .false. ! time-stepping: split-explicit (T) or implicit (F) time stepping
rn_zdfexp = 3 ! number of sub-timestep for ln_zdfexp-T
/
namdf_ric ! richardson number dependent vertical diffusion ("key_zdfrc")
rn_avmi = 100.e-4 ! maximum value of the vertical viscosity
rn_alp = 5. ! coefficient of the parameterization
rn_ric = 2 ! coefficient of the parameterization
rn_skufc = 0.7 ! Factor in the Ekman depth Equation
rn_mldmin = 1.0 ! minimum allowable mixed-layer depth estimate (m)
rn_mldmax = 1000.0 ! maximum allowable mixed-layer depth estimate (m)
rn_vtmix = 10.0 ! vertical eddy viscosity coeff [m2/s] in the mixed-layer
rn_wmix = 10.0 ! vertical eddy diffusion coeff [m2/s] in the mixed-layer
ln_mldv = .true. ! Flag to use or not the mixed layer depth param.
/
namdf_zke ! turbulent eddy kinetic dependent vertical diffusion ("key_zdfzke")
!
rn_diff = 0.1 ! coef. for vertical eddy coef. (avt-rn_ediff*rn_sqrt(e))
rn_diff = 0.7 ! coef. of the Kolmogoroff dissipation
rn_diff = 67.83 ! coef. of the surface input of the (=67.83 suggested when ln_mx10=7)
rn_emin = 1.e-6 ! minimum value of the [m2/s]
rn_emin0 = 1.e-4 ! surface minimum value of the [m2/s]
rn_mx1 = 2 ! mixing length: = 0 bounded by the distance to surface and bottom
rn_mx1 = 1 ! bounded by the local vertical scale factor
rn_mx1 = 2 ! first vertical derivative of mixing length bounded by 1
rn_mx1 = 3 ! as = 2 with distinct dissipative an mixing length scale
rn_mx1 = 1 ! Prandtl number function of richarson number (=1, avt=pol(Ri)*avm) or not (=0, avt=avm)
rn_mx10 = .true. ! surface mixing length scale = F(wind stress) (T) or not (F)
rn_mx10 = 0.04 ! surface buoyancy length scale minimum value
rn_lc = .true. ! Langmuir cell parameterisation (Axell 2002)
rn_lc = 0.15 ! coef. associated to Langmuir cells
rn_etau = 0 ! penetration of the below the mixed layer (ML) due to internal & inertial waves
rn_etau = 0 ! = 0 no penetration
rn_etau = 1 ! add a tke source below the ML
rn_etau = 2 ! add a tke source just at the base of the ML
rn_etau = 3 ! as = 1 applied on HF part of the stress ("key_coupled")
rn_etau = 0.05 ! fraction of surface tke value which penetrates below the ML (rn_etau=1 or 2)
rn_etau = 1 ! type of exponential decrease of the penetration below the ML
rn_etau = 0 ! constant 10 m length scale
rn_etau = 0.5m at the equator or to 30m poleward of 40 degrees
!
! ===== K-Profile Parameterization dependent vertical mixing ("key_zdkpp"), and optionally: "key_kppout" or "key_kplktk"
namdf_kpp ! K-Profile Parameterization dependent vertical mixing
!
ln_kpprnmix = .true. ! shear instability mixing
rn_difrmiv = 1.e-04 ! constant internal wave viscosity [m2/s]
rn_difisiv = 0.1e-04 ! constant internal wave diffusivity [m2/s]
rn_xriinfy = 0.8 ! local Richardson Number limit for shear instability
rn_difri = 0.0050 ! maximum shear mixing at Rig = 0 [m2/s]
rn_bvsqcon = -0.01e-07 ! Brunt-Vaisala squared for maximum convection [1/s^2]
rn_difcon = 1. ! maximum mixing in interior convection [m2/s]
rn_zvb = 0 ! horizontal averaged (-1) or not (=0) on avt and amv
rn_ave = 1 ! constant (=0) or profile (=1) background on avt
/
namdf_gls ! GLS vertical diffusion ("key_zdfgls")
!
rn_emin = 1.e-6 ! minimum value of e [m2/s]
rn_espmin = 1.e-12 ! minimum value of eps [m2/s]
ln_length_lim = .true. ! limit on the dissipation rate under stable stratification (Galperin et al., 1988)
rn_clim_galp = 0.53 ! Galperin limit
ln_urban = .true. ! Use Craig & Banner (1994) surface wave mixing parametrisation
ln_sigrpsi = .true. ! Activate or not Burchard 2001 mods on psi schmidt number in the vb case
rn_crhan = 100. ! Craig and Banner 1994 constant for vb flux
rn_charm = 70000. ! Charnock constant for vb induced roughness length
rn_tkebc_surf = 1 ! surface tke condition (0.1/2*pi*Neum/Dir Mellor-Blumberg)
rn_tkebc_bot = 1 ! bottom tke condition (0.1*pi*Neum)
rn_psibc_surf = 1 ! surface psi condition (0.1/2*pi*Neum/Dir Mellor-Blumberg)
rn_psibc_bot = 1 ! bottom psi condition (0.1*pi*Neum)

```

```

mn_stab_func = 2      ! stability function (0=Galf, 1=KG94, 2=GanotA, 3=GanotB)
mn_clos      = 1      ! predefined closure type (0=MY82, 1=k-eps, 2=k-w, 3=gen)
/
! namzdf_ddm   ! double diffusive mixing parameterization ("key_zdfddm")
rn_avrs     = 1.e-4   ! maximum avs (vertical mixing on salinity)
rn_hsbfr    = 1.6     ! heat/salt buoyancy flux ratio
/
! namzdf_tmrx ! tidal mixing parameterization ("key_zdftmrx")
rn_htmx     = 500.    ! vertical decay scale for turbulence (meters)
rn_n2min    = 1.e-8   ! threshold of the Brunt-Vaisala frequency (s-1)
rn_ne       = 0.333   ! tidal dissipation efficiency
rn_tmrx_itf = 0.2     ! ITF specific parameterisation
rn_tfe_itf  = .false. ! ITF tidal dissipation efficiency
/
! namzdf_tmrx ! tidal mixing parameterization ("key_zdftmrx")
rn_tmrx_itf = 1.      ! ITF tidal dissipation efficiency
/
! nammp        ! Massively Parallel Processing ("key_mpmp_mpi")
namct1      ! namct1          Massively Parallel Processing & Benchmark
namol       ! namol           Control prints & Benchmark
namolsol    ! namolsol         elliptic solver / island / free surface
/
! namolsol    ! elliptic solver / island / free surface
rn_solv     = 1      ! elliptic solver: =1 preconditioned conjugate gradient (pcg)
                   ! =2 successive-over-relaxation (sor)
mn_sol_arp  = 0      ! absolute/relative (0/1) precision convergence test
rn_rops     = 1.e-6   ! absolute precision of the solver
mn_rmin    = 210     ! minimum of iterations for the SOR solver
rn_rmax    = 800     ! maximum of iterations for the SOR solver
rn_rmod    = 10      ! frequency of test for the SOR solver
rn_remax   = 1.e-10  ! absolute precision for the SOR solver
rn_sor     = 1.96    ! optimal coefficient for SOR solver (to be adjusted with the domain)
/
! nammp        ! Massively Parallel Processing ("key_mpmp_mpi")
cn_mpi_send = 1,      ! mpi send/receive type ='S', 'B', or 'I', for standard send,
                   ! buffer blocking send or immediate non-blocking sends, resp.
mn_buffer   = 0      ! size in bytes of exported buffer ('B' case), 0 no exportation
ln_mngather = .false. ! activate code to avoid mpi_allgather use at the northfield
jmpi       = 4      ! jmpi number of processors following i (set automatically if < 1)
jpnj       = 8      ! jpnj number of processors following j (set automatically if < 1)
jpnij      = 32     ! jpni number of local domains (set automatically if < 1)
/
! namctl      ! Control prints & Benchmark
ln_ctl      = .true.  ! trends control print (expensive!)
mn_print   = 0       ! level of print (0 no extra print)
mn_iclts   = 0       ! start i indice of control sum (use to compare mono versus
                     ! multi processor runs over a subdomain)
mn_jclts   = 0       ! end i indice of control
mn_jctle   = 0       ! start j indice of control
mn_ispli   = 1       ! end j indice of control
mn_ispli   = 1       ! number of processors in i-direction
mn_ispli   = 1       ! number of processors in j-direction
mn_bench   = 0       ! Bench mode (1/0): CAUTION use zero except for bench
                     ! (no physical validity of the results)
mn_timing  = 1       ! timing by routine activated (-1) creates timing.output file, or not (-e)
/
! namr4        ! netcdf4 chunking and compression settings
("key_netcdf4")
namrd      ! namrd           dynamics and/or tracer trends
namflr    ! namflr          float parameters
namptr    ! namptr          Poleward Transport Diagnostics
namusb    ! namusb          Heat and salt budgets
/
! namc4        ! netcdf4 chunking and compression settings
("key_netcdf4")
namrd      ! namrd           dynamics and/or tracer trends
namflr    ! namflr          float parameters
namptr    ! namptr          Poleward Transport Diagnostics
namusb    ! namusb          Heat and salt budgets
/
! namcd4      ! netcdf4 chunking and compression settings
("key_netcdf4")
mn_nchunks_i= 4      ! number of chunks in i-dimension
mn_nchunks_j= 4      ! number of chunks in j-dimension
mn_nchunks_k= 31     ! number of chunks in k-dimension
mn_nchunks_k = jpk   ! setting mn_nchunks_k = jpk will give a chunk size of 1 in the vertical which
                     ! is optimal for postprocessing which works exclusively with horizontal slabs
ln_nczip   = .false.  ! (T) use netcdf4 chunking and compression
/
! namrdrd     ! diagnostics on dynamics and/or tracer trends
("key_trdyn" and/or "key_trdra")
("key_trmd1" or "key_trdrw")
ln_nczip   = .false.  ! (T) ignore chunking information and produce netcdf3-compatible files
/
! namrdrd     ! diagnostics on dynamics and/or tracer trends
("key_trdyn" and/or "key_trdra")
("key_trmd1" or "key_trdrw")
mn_trd      = 365     ! time step frequency dynamics and tracers trends
mn_tcls   = 0       ! control surface type in mixed-layer trends (0,1 or n_jjk)
mn_ucf    = 1.       ! unit conversion factor (-1 -> /seconds ; -86400. -> /day)
cn_trdrst_in = "restart_mid" ! suffix of ocean restart name (input)
cn_trdrst_out = "restart_mid" ! suffix of ocean restart name (output)
ln_trmid_restart = .false. ! restart for ML diagnostics
ln_trmid_instant = .false. ! flag to diagnose trends of instantaneous or mean ML T/S
/
! namflo      ! float parameters
("key_flo")

```



